Attenuation Correction of Reflectivity for X-Band Dual Polarization Radar

Hui Xiao¹, Yuxiang He^{1,2}, Daren Lü¹

1 Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing 100029 2 Department of Atmospheric Sciences, University of Illinois at Urbana-Champaign, IL, 61801

Abstract

X-Band Radar has been more and more widely used in rainfall estimation now, while its attenuation is more serious than S and C-Band radar. This constitutes a major source of error for radar rainfall estimation, in particular for intense precipitation. Thus, the X-Band dual polarization Doppler radar data needs to be corrected before used. On the basis of previous studies, this paper proposed a correction method which uses the characteristics of meteorology object and specific differential propagation phase $\sinh(K_{DP})$. Copolar differential phase shift is composed of two components, namely, differential propagation phase (Φ_{DP}) and back scatter differential phase (δ) . To estimate specific differential phase (K_{DP}) , at high frequencies, such as X-band, the angle (δ) is quite large, so these two components must be separated firstly at this frequency. The paper introduced an Optimal Recursive Data Processing Algorithm—Kalman filer which can separate these two components and filter out other random noise. Our experience indicated this method to be highly effective and practical and can also be used to process other radar signals. It is also very stable that the results of the reflectivity attenuation corrected using Kalman filter method processed K_{DP} .

Key Words Dual Polarization Radar, Attenuation Correction of Radar Reflectivity, Kalman filter

Corresponding author address: Dr. Yuxiang He, Department of Atmospheric Sciences, University of Illinois at Urbana-Champaign, 2204.Griffith Dr, Campaign, IL 61820. Email: heyx@illinois.edu; heyuxiangliu@gmail.com.

1 Introduction

It is one of the important world-class researches for precipitation radar to improve precision of area precipitation distribution quantitative estimation. Radar based on linear polarization is more and more widely used to do observation since Seliga and Bringi(1976) proposed to introduce polarization radar concept. Polarization technique of dual-linear polarization Radar is now of vital importance in remote detecting cloud precipitation physics. The polarization radar parameters (including differential propagation phase shift and differential reflectivity) which derived in the linear polarization basis are now widely used in precipitation estimation improvement (Qin et al., 2006). In Radar precipitation estimation, reflectivity attenuation is a key factor affecting the precision, thus must be corrected.

There were many studies in relation to S (10cm wave-length) and C (5cm wave-length) band radars. Besides, there are many countries use this two bands as operational radars, such as ChinaWSR-98D (S-band Weather Surveillance Radar-1998 Doppler) and USA WSR-88D (WSR-1988 Doppler). Since attenuation is relatively small for longer wave length (such as S) Radars, they have significant superiority detecting middle to heavy rain. For C band Radar, there are many studies using polarization and non-polarization method to do attenuation correction partly (Hildebrand, 1978). X-band Radar is now more and more widely used in field of research and operational application, the error induced by attenuation is very serious when using $Z_e - R$ (Z_e is equivalent radar reflectivity; R is rainfall intensity) to estimate precipitation. To solve the problem, X-band radar introduces polarization decomposition, by measure specific differential propagation phase shift K_{DP} to estimate precipitation which is more accurate than $Z_e - R$ relation. But, to a certain extent it is "noise" estimation, especially for light rain (Ryzhkov et al., 1996), $Z_e - R$ relation still has its superiority though attenuation heavily hindered its application. Thus study X-band Radar reflectivity attenuation correction then apply $Z_e - R$ to estimate precipitation is of great theoretical significance and practical value.

The main task for Reflectivity (Z_H) attenuation correction is to determine the relationship between attenuation rate (A_H) and distance. Once the relationship is determined, it is easy to correct Z_H for a specific distance. For routine single polarization Radar, A_H can be calculated indirectly by empirical relationship of $Z_H - R$ and $A_H - R$ (or $Z_H - A_H$) (Hildebrand, 1978). While the empirical relationship itself is very unsteady and sensitive to parameters such as system gain factor (Johnson et al., 1987). This is resulted from that A_H is derived by Z_H and Z_H itself has been attenuated. By theoretical analysis, Gorgucci et al (1998) showed that even without system correction error, the indirect utilization of attenuated reflectivity Z_H to do correction will introduce big errors.

While for dual polarization Radar, differential propagation phase shift Φ_{DP} can be used for reflectivity, which is more stable, correction. There were algorithms using Φ_{DP} for attenuation

correction (Bringi et al., 2002) most of which are suitable for S and C band Radar, little research has been conducted in applying the algorithms in X-band dual-polarization Radar (Anagnostou et al., 2004).

To acquire true Φ_{DP} value, various noises cancellation is the key process. If backscatter has following diagonal characteristics showed in Eq.[1] (Bringi et al., 2002), then differential propagation phase Φ_{DP} can be calculated from Mueller algorithm (Hubbert et al., 1995) showed by Eq. [2].

$$S_{BSA} = \begin{pmatrix} S_{hh} & 0 \\ 0 & S_{vv} \end{pmatrix}_{BSA} = e^{j\delta_{hh}} \begin{pmatrix} \left| S_{hh} \right| & 0 \\ 0 & \left| S_{vv} \right| e^{j(\delta_{vv} - \delta_{hh})} \end{pmatrix}_{BSA}$$
[1]

Where S_{BSA} is backscatter, S_{hh}, S_{vv} are horizontal and vertical scatter respectively, δ_{hh}, δ_{vv} are horizontal and vertical phase angle respectively.

$$\Phi_{DP} = \arg(\langle S_{vv} S_{hh}^* \rangle) + 2(k_h^r - k_v^i) r = \delta + 2K_{DP} r = \delta + \phi_{DP}$$
 [2]

Where δ is back scatter differential phase, $\phi_{\rm DP}$ is forward differential propagation phase shift. For Rayleigh scattering, δ is very small and can be ignored, while there are possibility high δ for non-Rayleigh scattering δ (Bringi et al., 2002) which is called δ effect.

Initial calculation shows that Rayleigh scatters are very small for light to middle rain, thus δ can be neglected. While for relatively strong rain areas, there may have serious deviations from Rayleigh scatter. Since forward and backward scattering are all include in differential phase shift Φ_{DP} observed by polarization decomposed Radar, it is very important and also very difficult to distinguish the two variables. If assume ϕ_{DP} as signal, then δ is the noise and it increases with raindrop size. Thus separate backward propagation phase shift to reduce this effect when using X-band dual-polarization Radar has vital effect on measurement results. Besides, meteorological object ambient disturb during detection, meteorological object own disturb as well as Radar machine interior noise all interfere differential propagation phase shift Φ_{DP} , make it fluctuating. And when using Φ_{DP} to do attenuation correction, only estimate the true value of forward differential propagation phase shift ϕ_{DP} accurately can use it to do attenuation correction.

According to characteristics of differential propagation phase shift, the paper introduces Kalman filter method, derives mathematical expressions for attenuation correction using differential propagation phase shift, and use the same method to cancel Radar signal high frequency noise and backward propagation phase shift δ . Then on the light of characteristics of stratiform cloud, obtains expression coefficients by fitting then do attenuation correction experiments for stratiform cloud.

2 Radar data filtration

The paper uses X-band dual-polarization Doppler radar of Institute of Atmospheric Sciences, Chinese Academy of Sciences (IAP-714XDP-A), and the performance parameters are listed in Table 1. The Radar adopts simultaneous dual transmit and receive (STAR) system in which every channel receives amplitude for every pulse:

$$\begin{bmatrix} s_{hh}^1 : s_{vv}^1 \end{bmatrix} \begin{bmatrix} s_{hh}^2 : s_{vv}^2 \end{bmatrix} \begin{bmatrix} s_{hh}^3 : s_{vv}^3 \end{bmatrix} \cdots \begin{bmatrix} s_{hh}^M : s_{vv}^M \end{bmatrix}$$
[3]

Where s_{hh} , s_{vv} are received signal by horizontal and vertical channel. Every channel at the same time has cross-polarization which is small compared to common-polarization and can be neglected. In order not to affect the precision of polarization measurement (such as Z_{DR}) during procession, we do not choose ground cluster filtering out scheme. The processor use following equation to calculated zero delay correlation coefficient:

$$\rho_{hv}(0) = \frac{\left\langle s_{vv} s_{hh}^* \right\rangle}{\sqrt{s_{hh} s_{vv}}}$$
 [4]

Including argument and phase:

$$\rho_{HV} = \left| \rho_{hv}(0) \right|$$

$$\Phi_{\mathit{DP}} = \arg \big[\rho_{\mathit{hv}}(0) \big]$$

Argument represents correlation coefficient, phase represents differential propagation phase shift.

Table 1 Performance parameters of the IAP-714XDP-A dual-polarization

Doppler radar with two channels

Radar characteristics variables	Detail description
Transmit system	Radar wave-length is 3.2cm, Total transmitting power ≥75KW, Single transmitting power≥ 35KW, Selectable pulse width (0.5µs, 1.0µs, 2.0µs)
Polarization diversity	Transmit horizontal and vertical polarized waves simultaneously
Antenna system	2.4m antenna diameter, 1.0° 3-dB beam width, 20°/second antenna speed
Antenna control system	PPI, RHI,VOL (volume scan) and sector scan modes, can customize parameters such as azimuth, elevation angle etc; solar calibration
Radar measurements	Horizontal polarization reflectivity, Doppler velocity, Spectrum width, Differential propagation phase shift, Ratio phase difference, differential reflectivity, correlation coefficient
Radar calibration	Calibrate antenna gain using signal generator, calibrate Radar position and orientation using solar method and GPS

2.1 Kalman filter

As long as reflectivity signal intensity strong enough not completely covered by noise, its phase can be measured. Under the condition of signal to noise rate *SNR* > 1, phase measurement always has superiority over power. Propagation phase shift is acquired through phase measurement which has numerous benefits comparing argument (such as intensity) measurement (Zrnic et al., 2006): 1) it is independent from calibration of transmitter and receiver; 2) not affected by attenuation; 3) not affected by beam fill degree; 4) not affected by ground cluster; 5) deviating from normal condition are rare even exists hailstone; 6) insensitive to droplet spectrum

variation; 7) it is easy to distinguish anomaly propagation.

Due to differential propagation phase shift Φ_{DP} has obvious fluctuations and belong to high frequency noise as for frequency domain. Thus many researchers (such as Hubbert et al., 1995) designed low-pass filter to remove the high frequency noises to keep the average trend of curve. But a specific filter relies on sample distance and the required smoothness (Hubbert et al., 1995). The paper utilize Kalman filter (Kalman, 1960) process Radar detected variable Φ_{DP} to acquire differential propagation phase shift Φ_{DP} needed by attenuation correction. Kalman filter uses mean square error as the optimal estimation criterion, seeks a set of recursive algorithm. It adopts space model of signal and noise state, using estimated value of previous time and observed value of current time to updates state variable, obtains current time estimation. In Kalman filter, visual physical meaning of time region language is adopted, that is only observational data during a limited time span are needed, thus relatively simple recursive algorithm can be applied and can be extended to unsteady random processes, and the data storage is small. Therefore, Kalman is suitable for computer real-time processing and calculating, it is an "Optimal Recursive Data Processing Algorithm".

When Radar remote sensing atmosphere, the to be measured state variable in every effective irradiation volume affected not only by atmosphere turbulence and other possible noises, but also by observational machine self noises. This can be abstracted to state variable estimation in discrete time process which can be described as the following discrete random differential equation

$$x_k = Ax_{k-1} + B(u_{k-1} + w_{k-1}), [5]$$

Define observation variable as z, the observation equation can be written as:

$$z_k = Cx_k + v_k \, \mathrm{d} \tag{6}$$

Table 2 shows meaning of each term in the model. w_{k-1} is random disturbance signals from atmosphere state (process noise), v_k is instrument noise (measurement noise), here assume they are independent from each other and the distributions are known.

Table 2 Meaning of each term of discrete difference equation

Symbols	Meaning
\mathcal{X}_k	System state
A	System matrix
B, u_k	State control variables
z_k	Measured value
C	Observational matrix
W_k	Procession noise $p(w) \sim N(0,Q)$
v_k	Measured noise $p(v) \sim N(0, R)$

Assume under ideal condition, there is a precipitation cell on Radar ray, its corresponding

relationship with Φ_{DP} is showed in Fig 1a. Ray enter precipitation cell from r_0 , Φ_{DP} increases. When the ray reaches cell center, Φ_{DP} change rate is greatest. Then ray passes cell center, Φ_{DP} increase rate slows down. After ray passes cell and enters non-precipitation area, Φ_{DP} stop increasing. If precipitation cell is homogenous, the differential propagation phase shift is indicated in Fig 1b, that is in the whole homogenous precipitation cell from r_0 to $r_{\rm m}$, Φ_{DP} increases with a constant rate, its slope indicates cell intensity, the more intense the precipitation cell, the greater the Φ_{DP} slope.

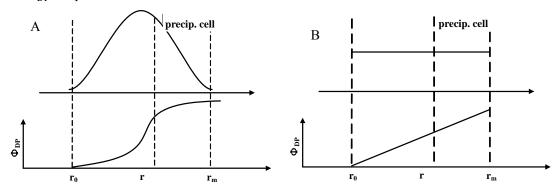


Fig. 1 Relationship between ideal precipitation cell and differential propagation phase shift Φ_{DP} (A. inhomogeneous precipitation cell, B. homogeneous precipitation cell)

During the procession, assume Radar ray propagation as the movement of moving object, the mathematic model for its kinematic equations can be described by following differential equation, thus describe the low of motion for differential propagation phase shift Φ_{DP} :

$$\begin{cases}
\Phi_{DP}(\mathbf{r}+1) = \Phi_{DP} + \mathbf{r}\Phi'_{DP}(\mathbf{r}) + \frac{1}{2}a(\mathbf{r})r^{2} \\
\Phi'_{DP}(\mathbf{r}+1) = \Phi'_{DP}(\mathbf{r}) + ra(\mathbf{r})
\end{cases}$$
[7]

Where $\Phi_{DP}(\mathbf{r})$ and $\Phi'_{DP}(\mathbf{r})$ are "location" and "velocity" of distance r respectively, here refer to accumulated and unit (K_{DP}) differential propagation phase shift respectively. Due to short average distance resulted from Radar adjacent distance base processing, further process must be do to use $\Phi'_{DP}(\mathbf{r})$ in Eq. [7]; $a(\mathbf{r})$ represents acceleration caused by uncertain factors such as meteorological object, ambient, Radar detecting system etc from r to r+1, it reflects undetectable caused by weather system itself and Radar detection system. $a(\mathbf{r})$ can be assumed to be a steady random sequence normally distributed which average value is zero and square error is Q, besides a(k) and $a(l)(l \neq k)$ are not related to each other, that is the expectation for $a(\mathbf{r})$ is $E\left\{a(\mathbf{r})\right\} = 0$, $E\left\{a(k)a(l)\right\} = Q\delta_K(k-1)$, where δ_K is Kronecker function whose characteristics are $\delta_K(k) = \begin{cases} 1, k=0 \\ 0, k \neq 0 \end{cases}$.

According to above discussion, Eqs. [5] and [6] change to:

$$X(r) = AX(r-1) + Ba(r-1)$$
 [8]

$$Z(\mathbf{r}) = CX(\mathbf{r}) + V(\mathbf{r})$$
[9]

Where,
$$X(\mathbf{r}) = \begin{bmatrix} \Phi_{DP}(\mathbf{r}) \\ \Phi'_{DP}(\mathbf{r}) \end{bmatrix}$$
, is state variable for observed weather system; $A = \begin{bmatrix} 1 & r \\ 0 & r \end{bmatrix}$, is

state transferring matrix for observed weather system; $B = \begin{bmatrix} \frac{r^2}{2} \\ r \end{bmatrix}$, is process noise, refer to the

corresponding weather system change noises when beam transfers from one irradiation volume to another during Radar detection; $C = \begin{bmatrix} 1 & 0 \end{bmatrix}$, is system observation matrix, Since differential propagation phase shift is directly observed, the first element in C matrix is equal to 1, $a(\mathbf{r})$ is Gaussian white noise sequence called dynamic white noise vector whose average value is zero and square error; $V(\mathbf{r})$ Gaussian white noise sequence called observational white noise vector whose average value is zero and square error, and does not correlated with $a(\mathbf{r})$; Known system state and observation equations, calculation can be accomplished through calculating prior state estimation value and prior error covariance matrix, modifying matrix, updating observation value, updating error covariance matrix etc (Zhang et al, 2001).

2.2 Filtration initialization

When using Kalman filtration algorithm, filtration initialization estimation value and error covariance matrix are needed to be assigned. Fig 3 shows that there are areas with great differential reflectivity fluctuation around Radar station during real Radar detection, which may induced by processing method of Radar signal processing system, machine interior noises and ground object noises etc. The fluctuation is varied up and down centered at a constant value. Bringi et al. (2002) indicates that Radar recorded differential phase shift is the result of Radar system and meteorological object, that is:

$$\arg(\rho_{co}^{meas}) = \arg(\rho_{co}) + \arg \iint e^{-j2(\Phi_h - \Phi_r)} f_h f_v d\Omega$$

$$-\left[(\Phi_{sw(h)}^t - \Phi_{sw(v)}^t) + (\Phi_{sw(h)}^r - \Phi_{sw(v)}^r) \right]$$

$$= \Phi_{DP} - (\phi_{DP})_{system}$$
[10]

Where Φ_{DP} is meteorological object differential phase detected by Radar; $(\phi_{DP})_{system}$ is the system differential phase shift which is a system error determined by the steady of system receiver and transmitter. Comparisons of different time data show that differential phase shift of Radar zero distance always drift to some extent which is not a constant value. This is related with Radar stability, system stability and fixed phase difference may be change with time and are controlled by temperature and humidity etc (Bringi et al., 2002). Therefore the first step in calculation is to determine initial phase that is system phase difference for every scan line. Known from above discussion, generally Φ_{DP} fluctuation increases with decreases of signal to noise rate. Since all of

them are unbiased estimation, the points have obvious fluctuation can be substituted by the average value of preceding several bases of every radial data. The paper uses 10 preceding points, that is, if Radar observed with 150m base, then the corresponding average value is at distance of 1.5km. The initial condition need to be appointed when using Kalman filter algorithm. To assign initial condition quickly during every scan line's processing, the paper uses the detected value of two initial bases after moving average to establish initial estimation, that is:

$$\hat{X}(2|2) = [Z(2) \qquad (Z(2) - Z(1))/r]^{T}$$
[11]

Initial estimation error is:

$$\widetilde{X}(2|2) = X(2) - \widehat{X}(2|2)
= \begin{bmatrix} \Phi_{DP}(2) \\ \Phi'_{DP}(2) \end{bmatrix} - \begin{bmatrix} Z(2) \\ (Z(2) - Z(1)) / r \end{bmatrix}
= \begin{bmatrix} -V(2) \\ \Phi'_{DP}(2) - \frac{\Phi_{DP}(2) - \Phi_{DP}(1)}{r} - \frac{V(2) - V(1)}{r} \end{bmatrix}$$
[12]

From state Eq. [13]:

$$\frac{\Phi_{DP}(2) - \Phi_{DP}(1)}{r} = \Phi'_{DP}(1) + ra(1)/2$$
 [13]

And

$$\Phi'_{DP}(2) - \frac{\Phi_{DP}(2) - \Phi_{DP}(1)}{r} = \Phi'_{DP}(1) + ra(1) - \Phi'_{DP}(1) - ra(1) / 2 = ra(1) / 2$$
 [14]

Thus Eq. [8] can be written as:

$$\widetilde{X}(2|2) = \left[\frac{-V(2)}{ra(1)} + \frac{V(1) - V(2)}{r} \right]$$
 [15]

Covariance matrix for error estimation is:

$$P(2 \mid 2) = E\{\widetilde{X}(2 \mid 2)\widetilde{X}^{T}(2 \mid 2)\} = \begin{bmatrix} Q^{2} & Q^{2} / r \\ Q^{2} / r & R^{2}r^{2} / 4 + 2Q^{2} / r \end{bmatrix}$$
[16]

2.3 Comparison between Kalman and other filter

To verify merits of Kalman filter, assume Radar detection distance is 150km, base length is 150m. There are weather signals as showed in Fig.2a within detection distance. Various noises induced the fluctuations of Radar echo signal, assume the noises are Gaussian white noise. Add Gaussian white noise to original signal with signal to noise rate 10, the result after adding noise showed in Fig.2b. Then use median and Kalman filter to filter contaminated signal, the results showed in Fig.2c and 2d. The figures show that median and Kalman filter have similar filter effect, they both keep the main characteristics of original signal. For median filter, the key problem is the width of the filter window, different widths give different results in experiment, see Fig. 2c. The two curves indicate results from two filter widths respectively. This shows that for different

frequency weather signal, optimal filter width is different which make it very difficult to real-time determine optimal width. While Kalman filter is non-sensitive to this. In experiment, assume process 360 Radar rays in a PPI, the computer main frequency is 3.0GHz, CPU time spend for median filter is 1.484s while Kalman is 0.031s. The time difference mainly caused by median filter must sort in every given filter width which cost tremendous computer time, and for other low-pass filters, due to large computer time spend on mutual transformation between frequency and time regions, are not suitable for real-time process. While Kalman filter have near real-time characteristics thus cost much less computer time.

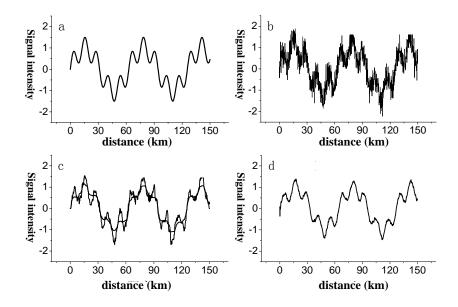


Fig2. Comparison of median and Kalman filter. (a) Orignal signal; (b) Orignal signal plus Gaussian noise; (c) Signal after median filtering; (c) Signal after by Kalman filtering. The ordinate means signal intensity and the abscissa is the distance from radar.

3 Attenuation correction

Since the very start of Radar utilization, there were many studies about attenuation caused by absorption and scatter when electromagnet passes precipitation region. Radar incident waves' attenuation is determined by attenuation cross section. Assume particles are ball shaped, incident wave length is λ , extinction coefficient σ_{ext} in Mie scattering area can be expressed as (Bringi et al., 2002):

$$\sigma_{ext} = \frac{2\pi}{k_0^2} \sum_{n=1}^{\infty} (2n+1) \operatorname{Re}(a_n^s + b_n^s).$$
 [17]

The order of $(D/\lambda)^5$ can be obtained if approximated by Mie scattering coefficient (a_1^s, b_1^s, a_2^s) , thus the extinction coefficient can be written as follows:

$$\sigma_{ext} = \frac{2\pi}{k_0^2} \left[3\operatorname{Re}(a_1^s + b_1^s) + 5\operatorname{Re}(a_2^s) \right]$$

$$= \frac{6\pi}{\lambda} \left(\frac{\pi}{6} D^3 \right) \left[i \left(\frac{\varepsilon_r - 1}{\varepsilon_r + 2} \right) \left\{ 1 + \left(\frac{\pi D}{\lambda} \right)^2 \left(t + u + \frac{5}{3} w \right) \right\} \right]$$
[18]

Where, \mathcal{E}_r is complex relative dielectric constant, definitions for t,u,w please refer to Hulst et al. (1981). Here Re represents real part of complex, $i=\sqrt{-1}$, D is diameter of particle. Bringi et al (2001) showed that following exponent relation can be used as first-order approximation for extinction coefficient:

$$\sigma_{\rm ext} = C_1 D^n \tag{19}$$

To use differential propagation phase shift do attenuation correction, the relation between unit differential propagation phase shift and attenuation rate must be solved first. According to Eq. [19], for 3cm wave length Radar, the particle's diameter is $0.1 \le D \le 8mm$. If use first-order approximation, n = 3.9, for particle's diameter $5 \le D \le 10mm$, the corresponding n = 4.6. To simplify discussion, the relation between attenuation and differential propagation phase shift is inducted, here $n \approx 4$. Thus attenuation rate can be expressed as:

$$\alpha_{h} = 4.343 \times 10^{3} \int \sigma_{ext}(D) N(D) dD; dB \ km^{-1}$$

$$\approx 4.343 \times 10^{3} C_{\lambda} \int D^{4} N(D) dD.$$
[20]

If use drop size distribution term to express K_{DP} , then:

$$K_{DP} = \frac{180\lambda}{\pi} \int \text{Re}[f_h(D) - f_v(D)] N(D) dD (\text{deg } km^{-1})$$
 [21]

Here f_h and f_v are channel forward scattering phase for horizontal and vertical polarization. Therefore attenuation correction based on unit differential propagation phase shift K_{DP} is (Park et al., 2005):

$$\alpha_h = c_h K_{DP} \tag{22}$$

Here a_h is attenuation rate which expresses effects from various factors affecting K_{DP} (Gorgucci et al., 2005). And the accumulative differential propagation phase shift Φ_{DP} is the accumulative variable of double K_{DP} . Then accumulative attenuation is:

$$A_{H}(r_{2}) - A_{H}(r_{1}) = 2 \int_{r_{1}}^{r_{2}} c_{h} K_{DP}(r) dr$$

$$= c_{h} [\Phi_{DP}(r_{2}) - \Phi_{DP}(r_{1})].$$
[23]

Here r_1, r_2 are two distances on Radar ray. If assign zero to r_1 , that is the first base for Radar, then the above equation changes to:

$$A_{H}(r) = C_{h} [\Phi_{DP}(r) - \Phi_{DP}(0)]$$
 [24]

Where, $A_H(r_1) = 0$, $\Phi_{DP}(0)$ is differential propagation phase shift for the first base of Radar ray, that is $(\phi_{DP})_{system}$ in Eq. [10], which represents system initial differential propagation phase shift. Then the attenuation correction for Radar reflectivity can be expressed as (Park et al., 2005):

$$Z(r) = Z_m(r) + A_H(r) = Z_m(r) + c_h [\Phi_{DP}(r) - \Phi_{DP}(0)]$$
 [25]

Here $Z_m(r)$ is Radar detection value at distance r; $A_H(r)$ is correction at distance r use differential propagation phase shift; Z(r) is the result after correction. It can be seen that they are linearly related when using differential propagation phase shift to correct reflectivity.

4 Case analysis

On May 15, 2007, a low trough controlled the east of China; the northeastern China locates in front of 500hPa trough. On 850hPa level, the trough stretches southward from the northeastern China till South of Yangtze River Basin at 08 o'clock (Beijing Time, same as below). At 14 o'clock, there a low vortex was generated above Jilin area which maintained at Jilin Province in the northeastern China till 20 o'clock. It can be seen that the low vortex cloud system developed and moved eastward from the May 15 satellite cloud map. The sounding profile at Changchun station shows that the difference between dew point and temperature is big with unsaturated lower atmosphere at 8 o'clock May 15. At 20 o'clock, the lower atmosphere became saturated. Around 21 O'clock May 15, there was precipitation in Yitong area, Jilin Province.

4.1 Case filtration

Fig.3 is the Radar ray with 195° azimuth and 1° elevation angle, the variance of Φ_{DP} with distance is showed by dash lines. Following processing is based on Radar polar coordinate system. The figure shows that differential propagation phase shift departure from theoretical value increases with distance accompanied by fluctuations. The fluctuations mainly induced by atmosphere self fluctuation and Radar machine interior detection noise etc, that is processing and measurement noises in Kalman filter. Fluctuations are obvious especially around Radar station and far part of Radar ray. According to above discussions, the fluctuations are unbiased and have big initial phase shit at the same time. In this case, the initial phase is 103°, and there appear fluctuation signals with long nail shape on Radar ray, the signal is an obvious backward propagation phase shift δ effect. Compared with reflectivity in Fig 5, we find that at 5km and 11km where δ effect appears, the reflectivity is greater than adjacent. Besides obvious δ effect, there also exists unobvious backward propagation effect which structure showed in Fig 5 as minor fluctuation, and high frequency noises from other noise source. The object of this paper using Kalman filter is to solve the problems and obtain differential propagation phase shift curve used by attenuation correction.

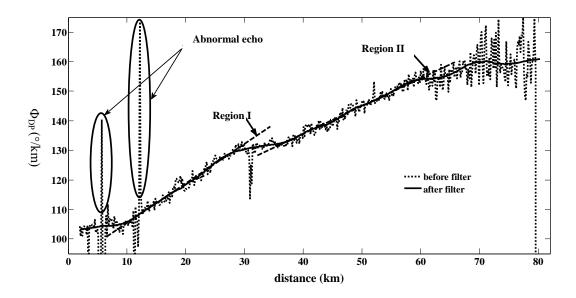


Fig.3 Differential propagation phase shift with distance before and after filtration. The ordinate means differential propagation phase shift $\Phi_{DP}(^{\circ}/\text{km})$ and the abscissa is the distance from radar.

During detection, Radar signal is contaminated by atmosphere unsteady noise and Radar system interior machine noise. Using filtered Radar ray showed as real line in Fig. 3, it can be seen that noise effect is reduced and the backward scattering δ effect of polarization Radar observed Φ_{DP} (abnormal echo in Fig 3) is also canceled

The differential propagation phase shifts of RHI data for the same time are processed use Kalman filter, the results showed in Fig 4. The figure shows that differential propagation phase shift Φ_{DP} obviously improved after filtration. The figure has great fluctuations before filtration which directly affect the stability and correct of afterward attenuation correction, while after filtration, the figure is more smooth basically filter out δ effect and random noises.

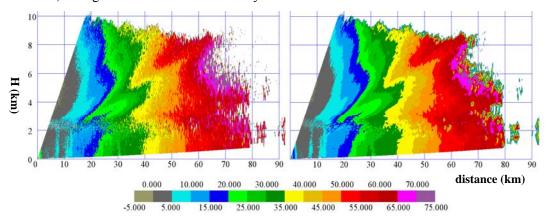


Fig.4 Comparison of differential propagation phase shift $\Phi_{DP}(^{\circ}/km)$ before and after filtration, The ordinate means height H (km) and the abscissa is the distance R (km) from radar.

4.2 Attenuation correction for a case

Use the weather event of May 15, 2007 as an example to show how to obtain correction parameter c_n by detecting real meteorological weather object. According to conceptual model in Fig 1, differential propagation phase shift varying linearly at 10-30kn and 30-60km, thus divide differential propagation phase shift to two regions as showed in Fig 3. Region I has greater differential propagation phase shift variation than Region II which results in stronger echo in Region I than in Region II. And the precipitation cells are homogenously varied in the two regions which indicate that the clouds in both regions are relatively homogeneous stratiform clouds. Followed use this constrained relation, by statistical regression to determine attenuation correction coefficient, then further determine the relation between attenuation rate and distance.

Use Eq. [21] do regression for the two regions to make corrected reflectivity consistent in the two regions, that is, make sure echo intensities are homogeneous in the two regions. The mathematical constraint is mean statistical square error. The regression corrected results are showed in Fig. 5.

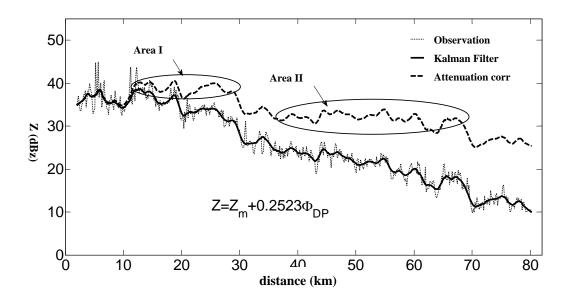


Fig.5 Comparison of radar reflectivity factor along a ray distance before and after correction. The ordinate means radar reflectivity factor Z (dBz) and the abscissa is the distance R (km) from radar. The thin dot line stands for the observation, Thick line is for the result with Kalman filter and thick dashed line is for the result after attenuation correction.

Comparison of phase shift and reflectivity curves of differential propagation phase shift shows that reflectivity curve attenuates obviously with distance before correction (differential propagation phase shift figure does not show such an obvious attenuation). Reflectivity generally consistent in regions I and II after correction, which coincides with detection conclusions of differential propagation phase shift. It can be seen that the radar system can use $Z = Z_m + c_h \Phi_{DP}$ conduct reflectivity attenuation correction, for our case $c_h = 0.2523$. The before mentioned Kalman filter is also used during reflectivity processing, short dash line in Fig. 5

represents observed value and solid line is results after filtration. Here filter reflectivity to a certain degree is good for understanding its main meteorological information. The method utilizing real echo to determine correction coefficient has great practical value and always more effective.

To verify above results in our radar system, apply the correction relation to correct radar detected RHI data, the results are showed in Fig. 6.

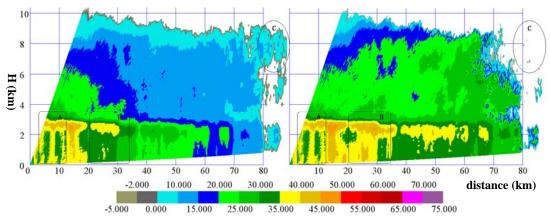


Fig.6 Comparison of radar reflectivity before and after correction. The ordinate means height H (km) and the abscissa is the distance R (km) from radar.

In general, during the process, zero-layer bright band has earthward trend with distance before correction, while basically keep horizontal after correction. It is a stratiform cloud precipitating, the atmosphere is stable and horizontally homogeneous. Thus the trend before correction is incorrect and coincides with weather situation after correction.

Comparison between before and after correction figures show that the correction method has numerous merits: The two figures hardly have differences before and after correction for A region in due to light attenuation since Radar ray does not pass strong precipitation area. Do statistical to reflectivity of each point with a constant interval ΔZ before and after correction for A region, the number distribution showed in Fig.8a and 8b. The figure shows that number distribution pattern keep unchanged before and after attenuation correction with average values 40.438dBz and 40.698dBz respectively which has little difference. Thus the basic characteristics keep unchanged in weak precipitation area after attenuation correction this way.

Radar ray appear obvious attenuation in B and further regions after passed strong precipitation region between regions A and B in Fig. 6. Following above statistical method, results for the B and further regions are showed in Fig.7c and 7d. The figures show that the average value is 27.184dBz and 37.321dBz which indicate serious attenuation for the entire B and further regions. At the same time, the reflectivity number distribution spectrum width narrowed after correction, while spectrum distribution pattern, such as distribution of peak numbers, basically keep unchanged. Such change of spectrum pattern further illustrates that the stability of the correction method. In Fig.7d, there appears many small echoes after correction between 5-25dBz

which not exist before correction. This is because this part of echo has already attenuated. These entirely attenuated echoes have signal to noise rate nearly 1. This kind of weak echo can be reconstructed as long as SNR > 1 when using differential propagation shift Φ_{DP} do attenuation correction. It is credible due to self characteristics of differential propagation phase shift. While it is incredible to correct this kind of echo through intensity superposition, since the echo is hard to restore and has great fluctuations. For region C in Fig. 6, great discrepancy showed before and after correction. There is echo before correction, while it disappeared after correction. The region may caused by tri-body echo induced by zero-layer light band. This kind of tri-body echo has small velocity, large and litter pulsation, small correlation coefficient detected by polarization Radar, ground object (such as wind blowing leaf etc) induced Doppler frequency shift can be positive or negative with great noise and fluctuated around minimum value. In Fig.4, there is no differential propagation phase shift value in the region but has echo on intensity image which indicates the false echo removal ability using the correction method.

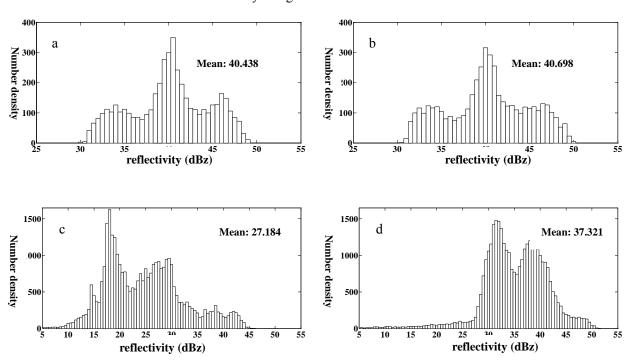


Fig. 7 Distributions of reflectivity number density for various regions before and after correction. a,b. Radar ray in front of strong precipitation region, c,d. Radar ray behind strong precipitation region, left to right column: before and after correction.

The ordinate means reflectivity number density and the abscissa is the reflectivity (dBz) from radar.

Fig.8 is isograms comparison before and after correction. It shows isograms attenuated rapidly with distance before correction, such as isograms of 25dBz, 20dBz, 15dBz above zero-layer light band, show arch distribution with distance. The isograms nearer obviously greater than the ones far from Radar Below zero-layer light band as showed in Fig.8a. The isograms do not show obvious change trend with distance after correction. Besides, zero-layer wider a little

with distance which mainly caused by beam width widen with distance, see Fig. 8b.

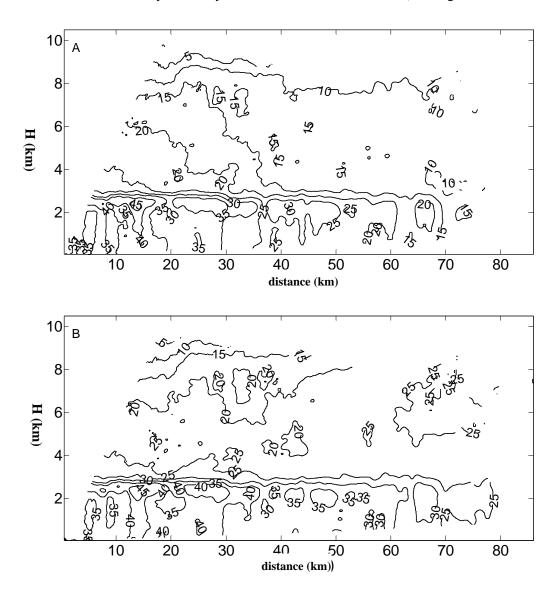


Fig.8 Isograms for radar reflectivity before and after correction (A. before correction; B. after correction). The ordinate means height H (km) and the abscissa is the distance R (km) from radar.

5 Summary and discussion

The most difficult problem using X-band Radar is its attenuation much greater than S or C-band. Radar reflectivity decreased obviously with the propagation of radar ray in precipitation media. The study introduced an attenuation correction method according to precipitation media self characteristics. The key problem when using differential propagation phase shift does attenuation correction is the quality control of differential propagation phase shift curve. Various noise sources are affecting real differential propagation phase shift detection, including backward phase shift δ effect, rising and falling of meteorological object, radar machine interior noises etc. All of these noises can be remove to some extent through filtration. The paper uses Kalman filter

to deal with differential propagation phase shift Φ_{DP} of dual-polarization radar. This method is an optimal self regression data process algorithm which makes it an optimal and most efficient for all problems solving. Each term has definite physical meaning and has great practical value. According to attenuation correction method, if we do not remove noises in Radar received differential propagation phase shift before correction, there is high possibility that backward propagation phase shift δ effect and other noises may much greater than differential propagation phase shift itself which made serious distortion to corrected reflectivity then exert huge negative effect on stability and accuracy of correction. The study according to the self characteristics of differential propagation phase shift, utilizes the observed data by the X-band dual polarization Doppler radar of Institute of Atmospheric Physics, Chinese Academy of Sciences, through stable stratiform cloud case study, compares changes of RHI images and number density distribution before and after correction, indicates good stability of filtered propagation phase shift to correction of reflectivity factor. For weak echo situation, the stability of the correction method is like Radar nearby area of stratiform cloud in the paper. Further investigation of many weak echo cases show that the stability and accuracy of the correction method only depend on accuracy of differential propagation phase shift not matter weak or strong echo. Under weak echo condition, though backward propagation phase shift δ is small, other noises still exists and more serious, result in necessary to filter differential propagation phase shift. It has certain practical value to determine attenuation correction coefficient statistically by collecting numerous homogeneous statiform cloud precipitation detection data of different time, same location, same Radar, on the basis of filtration. Further studies are needed to establish robust attenuation correction coefficient suitable for the Radar system by more observational data collection.

Acknowledgments

This work was supported by the National Natural Science Foundation of China (Grant No. 40875080) and Ministry of Science and Technology of China (Grant No. 2006BAC12B01-01).

References

- Anagnostou, E.N., M.N. Anagnostou, W.F., Krajewski, A. Kruger, and B.J. Miriovsky, 2004. High-Resolution Rainfall Estimation from X-Band Polarimetric Radar Measurements[J]. Journal of Hydrometeorology, 5(1): 110-128.
- Bringi, V.N., 2001. Correcting C-Band Radar Reflectivity and Differential Reflectivity Data for Rain Attenuation: A Self-Consistent Method With Constraints[J]. IEEE Transcations Geoscience and Remote Sensing, 39(9): 1906-1915.
- Bringi, V.N. and V. Chandrasekar, 2002. Polarimetric Doppler Weather Radar: Principles and Applications[M]. Cambridge University Press.
- Gorgucci, E. and V. Chandrasekar, 2005. Evaluation of Attenuation Correction Methodology for Dual-Polarization Radars: Application to X-Band Systems[J]. Journal of Atmospheric and Oceanic Technology, 22(8): 1195-1206.
- Gorgucci, E., G. Scarchilli, V. Chandrasekar, P.F. Meischner and M. Hagen, 1998. Intercomparison of Techniques to Correct for Attenuation of C-Band Weather Radar Signals[J]. Journal of Applied Meteorology, 37(8): 845-853.
- H. C.van de Hulst, 1981. Light Scattering by Small Particles[M]. New York: Dover.
- Hildebrand, P. H, 1978. Iterative Correction for Attenuation of 5 cm Radar in Rain[J]. Journal of Applied Meteorology, 17(4): 508-514.
- Hubbert, J. and V.N. Bringi, 1995. An Iterative Filtering Technique for the Analysis of Copolar Differential Phase and Dual-Frequency Radar Measurements[J]. Journal of Atmospheric and Oceanic Technology, 12(3): 643-648.
- Johnson, B.C. and E.A. Brandes, 1987. Attenuation of a 5-cm Wavelength Radar Signal in the Lahoma-Orienta Storms[J]. Journal of Atmospheric and Oceanic Technology, 4(3): 512-517.
- Kalman, R.E, 1960. A new approach to linear filtering and prediction problems[J]. Transactions of the ASME. Series D: Journal of Basic Engineering, 82:35-45.
- Park, S.G., V.N. Bringi, V. Chandrasekar, M. Maki, and K. Iwanami, 2005. Correction of Radar Reflectivity and Differential Reflectivity for Rain Attenuation at X Band. Part I: Theoretical and Empirical Basis[J]. Journal of Atmospheric and Oceanic Technology, 22(11): 1621-1632.
- Qin Y., B. Li and P. Zhang, 2006. A Study of the Relationship Between Radar Reflectivity of Rain and Relative Humidity of Atmosphere[J](in Chinese). Chinese Journal of Atmospheric Sciences, 30(2):351-359.
- Ryzhkov, A. and D. Zrni, 1996. Assessment of Rainfall Measurement That Uses Specific Differential Phase[J]. Journal of Applied Meteorology, 35(11): 2080-2090.
- Seliga, T.A. and V.N. Bringi, 1976. Potential Use of Radar Differential Reflectivity Measurements at Orthogonal Polarizations for Measuring Precipitation[J]. Journal of Applied Meteorology, 15(1): 69-76.
- Wiener, N, 1949. Extrapolation, Interpolation, and Smoothing of Stationary Time Series[J]. New York: Wiley.
- Zhang P., B. Du and T. Dai, 2001. Radar Meteorology[M](in Chinese), Beijing: China Meteorological Press,511.

Zrnic, D.S., V.M. Melnikov, and J.K. Carter, 2006. Calibrating Differential Reflectivity on the WSR-88D[J]. Journal of Atmospheric and Oceanic Technology, 23(7): 944-951.